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Oil-Mineral Flocculation and Settling Velocity in Saline Water

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Abstract

Cohesive particles in aquatic systems can play an important role in determining the eventual fate of spilled oil via the generation of Oil-Mineral Aggregates (OMAs). Series of laboratory experiments have been conducted aiming at filling the knowledge gap regarding how cohesive clay particles influence the accumulation of petroleum through forming different aggregate structures and their resulting settling velocity. OMAs have been successfully created in a stirring jar with artificial sea-water, crude oil and two types of most common natural cohesive minerals, Kaolinite and Bentonite clay. With the magnetic stirrer adjusted to 490 rpm to provide a high level homogeneous flow turbulence (Turbulence dissipation ϵ estimated to be about $0.02 \text{ m}^2 \cdot \text{s}^{-3}$), droplet OMAs and flake/solid OMAs have been obtained in oil-Kaolinite sample and oil-Bentonite sample, respectively. Kaolinite clay with relatively low flocculation rate ($R_f = 0.13 \text{ min}^{-1}$) tends to physically

25 attach around the surface of oil droplets. With the lower density of oil, these
 26 oil-Kaolinite droplet OMAs generally show lower settling velocity compar-
 27 ing to pure mineral Kaolinite flocs. Differently, Bentonite clay with higher
 28 flocculation rate ($R_f = 0.66 \text{ min}^{-1}$) produces high porous flocs that can ab-
 29 sorb or be absorbed by the oil and form compact flake/solid OMAs with
 30 higher density and settling velocity than pure Bentonite flocs. Some of the
 31 large oil-Bentonite OMAs are measured to have settling velocities up to 10
 32 $\text{mm}\cdot\text{s}^{-1}$. In the mixture condition (Kaolinite:Bentonite = 1:1 in weight), oil
 33 can be considered to preferably interacting with Bentonite. The microflocs
 34 ($< 160 \text{ }\mu\text{m}$) are dominated by Kaolinite with lower oil participation and
 35 lower settling velocity, but the Macroflocs ($> 160 \text{ }\mu\text{m}$) which dominates the
 36 entire sample's average properties are more influenced by the oil-Bentonite
 37 characteristics.

38 *Keywords:* Oil-Mineral Aggregates (OMAs), settling velocity, microfloc,
 39 Macrofloc.

40 1. Introduction

41 Since the petroleum exploration and transportation became one of the
 42 most critical industrial activities for the global economic growth, extremely
 43 large oil spill disasters, such as the 1989 Exxon Valdez spill (0.26 million
 44 barrels of oil released to Alaska's Prince William Sound) ([Peterson et al.,](#)
 45 [2003](#)) and the 2010 Deepwater Horizon (DWH) disaster (4.9 million barrels
 46 of oil were released into the Gulf of Mexico) ([Crone and Tolstoy, 2010](#); [At-](#)

las and Hazen, 2011), and increasingly smaller size spills occurred in the coastal zones (Hayakawa et al., 2006; Doshi et al., 2018; Liu and Callies, 2019). These oil spill accidents pose detrimental impacts on sea-based human activities (Peterson et al., 2003), ecosystem contamination of aquatic bio-communities (Ainsworth et al., 2018) such as fishes (Murawski et al., 2014), birds (Henkel et al., 2012), coral (White et al., 2012) or plankton (Almeda et al., 2013, 2016). Although most mitigation methods focus on spilt oil floating onto the water surface (Reddy et al., 2002, 2012; Liu et al., 2012), there can be a considerable portion of spilt oil settles to the sea-floor after flocculating with natural cohesive materials, including sediments and organic particles (Chanton et al., 2014; Yan et al., 2016; Jones and Bridgeman, 2016; Romero et al., 2017; O’Laughlin et al., 2017).

Flocculation with cohesive mineral sediments can be especially common in more energetic coastal environments where resuspension of sediment is more likely or near river mouths where new supplies of sediments are abundant (Strom and Keyvani, 2016; Shen et al., 2018). When crude oil is released into aquatic systems in nature, oil droplets can be open to flocculate with suspended particles (Sterling Jr et al., 2005). Through settling and deposition, the oil mineral aggregates may eventually preserving in the sea-floor depositions over geological time (Romero et al., 2017). Therefore, the interactions of oil and aquatic mineral particles, or biological materials can play an important role in the fate of spilt oil (Khelifa et al., 2002, 2005a; Passow and Hetland, 2016; O’Laughlin et al., 2017). This study focuses on

70 the influence of mineral sediments on oil droplets through flocculation. Some
71 insights into the flocculation of oil droplets with biological materials, namely
72 the marine snows, can be found in, for instance, [Passow et al. \(2012\)](#) and a
73 comprehensive review article of [Daly et al. \(2016\)](#).

74 Oil droplets tend to aggregate with, and finally be stabilized by, cohesive
75 particles or suspended particle materials (SPM) in the water column and
76 form oil-mineral aggregates (OMAs) ([Khelifa et al., 2002](#)), oil-SPM aggre-
77 gates (OSAs) ([Khelifa et al., 2005a](#)) or oil-particle aggregates (OPAs) (e.g.,
78 [Zhao et al. 2014, 2016, 2017](#)). In the present study, the term “OMAs” has
79 been used because only mineral clay has been utilized to flocculate with oil
80 droplets in the cases. Several earlier studies focus on the structure of OMAs
81 using microscopy imagery and “droplet OMAs”, “flake OMAs” or “solid
82 OMAs” are most commonly observed OMA structures ([Lee and Stoffyn-
83 Egli, 2001](#); [Stoffyn-Egli and Lee, 2002](#)). Droplet OMAs are combination or
84 enclosure of one or several oil droplet(s) and mineral particles/flocs via sur-
85 face attachments. On the contrary, flake OMAs have a similar shape as
86 solid OMAs which have membrane-like sheets with an orderly arranged oil
87 and mineral particle configuration ([Stoffyn-Egli and Lee, 2002](#)). With higher
88 shear strength, structures of flake OMAs could be altered to become solid
89 OMAs because the crumpling or breaking of flake type OMAs may form more
90 compact and denser floc structures ([Stoffyn-Egli and Lee, 2002](#); [Loh et al.,
91 2014](#)). They can be highly compact with oil and mineral absorbed together
92 and organized in dendritic or foldable feather-shape structures.

93 Later, several important OMAs studies have been focused on the for-
94 mation mechanisms and influence factors. [Omotoso et al. \(2002\)](#) presented
95 a flocculation index based on the sedimentation behavior of a sheared oil-
96 mineral-water mixture. It was used to quantify the degree of interaction of
97 oil and minerals in water which was found to be dependent on the viscosity
98 of the crude oil and the type of mineral present. On the other hand, [Khelifa](#)
99 [et al. \(2002\)](#) suggested that droplet shape and size were not correlated to
100 oil viscosity, but the concentration of oil droplets decreased rapidly with oil
101 viscosity, temperature and asphaltenes-resins content (ARC). [Le Floch et al.](#)
102 [\(2002\)](#) quantified the amount of oil incorporated into OMA with the salinity
103 ranging from 0~35 ppt. They demonstrated that the OMA formation was
104 significantly enhanced by salinity when comparing to distilled water condi-
105 tion. However, the amount of oil contained in OMAs saturated at low salinity
106 of only 2 ppt and further enhancing salinity showed almost no effect OMA
107 formation. This salinity threshold depends on other parameters including oil
108 type and the nature of the mineral present. Below this salinity threshold,
109 there is a linear decrease in the amount of oil incorporated in OMA, to prac-
110 tically zero in distilled water. [Hill et al. \(2002\)](#) presented an equation that
111 defines the time required to coat and stabilize oil droplets with mineral par-
112 ticles suspended in a turbulent medium. The finding that OMA form rapidly
113 given adequate sediment concentration should play a key role in oil spill re-
114 sponse decision. With the high demand of quantitative understanding in the
115 formation of OMAs processes, some further laboratory experimental studies

116 have been reported. [Khelifa et al. \(2005b\)](#) presented the laboratory results
117 showing a positive correlation of OMA sizes and the concentration of mineral-
118 stabilized droplets with salinity positively from zero to a critical aggregation
119 salinity in the range of 1.2~3.5 ppt. And it is believed that the effect if
120 salinity on droplet size distribution is strongly influenced by clay type. More
121 recently, [Sun et al. \(2010, 2013\)](#) presented the experimental results showing
122 that the formation of oil suspended particle matters aggregates increased
123 exponentially with the mixing time and reached an equilibrium within 4~5
124 hours at a provided turbulence dissipation rate of $2.6 \text{ m}^2\cdot\text{s}^{-3}$. They believed
125 the shaking rate (turbulence) largely influences the maximum oil trapping
126 efficiency in OMAs. And it is found that most of the formed aggregates were
127 solid aggregates and single droplet aggregates with low mixing energies, and
128 multi-droplet oil suspended particles aggregates with high mixing energies.

129 Among all the previous OMA literature, very limited studies have been
130 reported to systematically investigate OMAs settling dynamics. [Khelifa et al.](#)
131 [2008](#) reported a series of detailed laboratory experimental jar tests on chem-
132 ical dispersed oil and natural mixture sediment aggregation. Their data,
133 probably for the first time, showed a direct relationships between the mea-
134 sured settling velocity and OMA size. They suggested that those flocs with
135 low oil concentration may barely change the oil-sediment aggregates behav-
136 iors, but with high oil concentration within the oil-sediment aggregates, their
137 density can be much smaller than pure sediment flocs having similar floc sizes.
138 For most sediment types they tested, the effective density of the oil-sediment

139 aggregates can be about 2~3 times less than those of pure sediment flocs.
140 Importantly, in their samples, it is suggested that the presence of chemical
141 dispersed oil may enhance the stickiness of sediment grains which helps build-
142 ing up the large flocs with oil participation. More recently, [O’Laughlin et al.](#)
143 [\(2017\)](#) reported measured results of dilbit-derived OMAs settling velocity
144 from series of laboratory and wave tank experiments in response to the pres-
145 ence or absence chemical dispersants. They suggested that settling velocities
146 of artificially formed OMAs on the order of $0.1\sim0.4\text{ mm}\cdot\text{s}^{-1}$. Moreover, the
147 OMA size, settling velocity and effective particle density were increased in
148 response to the higher concentration of suspended sediment. Their data also
149 show evidences that dispersant may inhibits flocculation. These two stud-
150 ies clearly indicated the importance of cohesion(stickiness) in determining
151 the resulting oil-floc and their settling velocity. As mentioned above, earlier
152 studies also provided very comprehensive understanding on the structures
153 of OMAs. Here, we further hypothesize that a main factor controlling the
154 structures of OMA is the properties of the mineral sediments, for example,
155 their stickiness. Hence, a reasonable next step is to further relate different
156 type of OMA structures with their settling characteristics.

157 The present study is motivated to investigate the effect of mineral types,
158 which provide different stickiness, in determining the OMA structures and
159 the resulting settling velocities. Data obtained from the controlled labora-
160 tory experiments are analyzed with three main objectives: 1) To understand
161 the OMAs structures formed with different types of common clay minerals

162 by high-resolution digital microscopy, 2) to measure physical characteristics
163 of OMAs, such as their sizes and settling velocities using LabSFLOC-2 cam-
164 era, and the most importantly 3) synthesize measured data to gain insights
165 into OMA structure and settling dynamics due to different clay types. The
166 remaining of this paper is organized as follow. Section 2 focuses on the labo-
167 ratory methods, including the OMA generation, turbulence characterization,
168 microscopy and LabSFLOC-2 system for studying settling characteristics.
169 Results are presented in terms of OMA structures and characteristics in Sec-
170 tion 3, discussions are in Section 4 and important concluding remarks are in
171 Section 5.

172 2. Materials and methods

173 2.1. Laboratory experiment setup

174 An experimental stand set (Figure 1a) has been designed and a series
175 of magnetic stirring jar experiments have been conducted at the Center
176 for Applied Coastal Research, University of Delaware. White Kaolin clay
177 (92.3 ± 2.5 % Kaolinite), Wyoming sodium Bentonite clay (85.2 ± 2.3 % Mont-
178 morillonite) (two most common clay types with large difference in cohesion
179 in saline water) and raw Texas crude oil (Dynamic viscosity: 7.27×10^{-3} Pa·s
180 at 20 °C) with various proportions are used to generate OMAs. As summa-
181 rized in Table 1, we specify oil-to-sediment ratio close to 2 with clay mineral
182 concentration of 0.5 g per litre of saline water, which provide a condition
183 for maximum OMA formation efficiency condition according to the previ-

ous studies ([Guyomarch et al., 2002](#); [Khelifa et al., 2008](#); [Ajjolaiya et al., 2006](#)). Artificial seawater (Salinity ≈ 35 ppt) has been made from mixing clean water and pure salt. The jar has a diameter of 11 cm and the flow depth is 13 cm (1 liter salt water). Magnetic stirring speed is set to 490 rpm (Device range: 0~1000 rpm) for providing the constant turbulence intensity for OMAs generation. Three-component flow velocities are measured by a Vectrino Profiler (Nortek), which was mounted on the shelf above the magnetic stirrer with the sensor probes located 5 cm below the water surface in the jar (in Figure 1a). Flow velocity data was collected without crude oil and sediment but in otherwise the same flow conditions (artificial seawater in the jar with same flow depth). The time series of turbulent velocity fluctuations are transformed into Fourier space to obtain turbulent kinetic energy spectrum. Turbulence dissipation rate is then estimated to be $\epsilon \approx 0.02 \text{ m}^2\cdot\text{s}^{-3}$ via matching the Kolmogorov spectrum with Taylor frozen turbulence approximation (e.g., [Voulgaris and Trowbridge 1998](#); [Huang et al. 2018](#)). The corresponding shear parameter is of seawater (Salinity ≈ 35 ppt) at 20 °C (Viscosity $\approx 1.08 \times 10^{-3} \text{ Pa}\cdot\text{s}$).

Different types of mineral flocs and Oil-Mineral Aggregates (OMAs) samples are generated, including 1) Kaolinite flocs, 2) Bentonite flocs, 3) Mixed Kaolinite-Bentonite flocs, 4) Oil-Kaolinite aggregates, 5) Oil-Bentonite aggregates and 6) Oil-Kaolin-Bentonite aggregates. Each experimental run last up to 2 hours and OMAs are allowed to settle down overnight (~ 8 hours) which should be long enough for all the particles aggregation and settling.

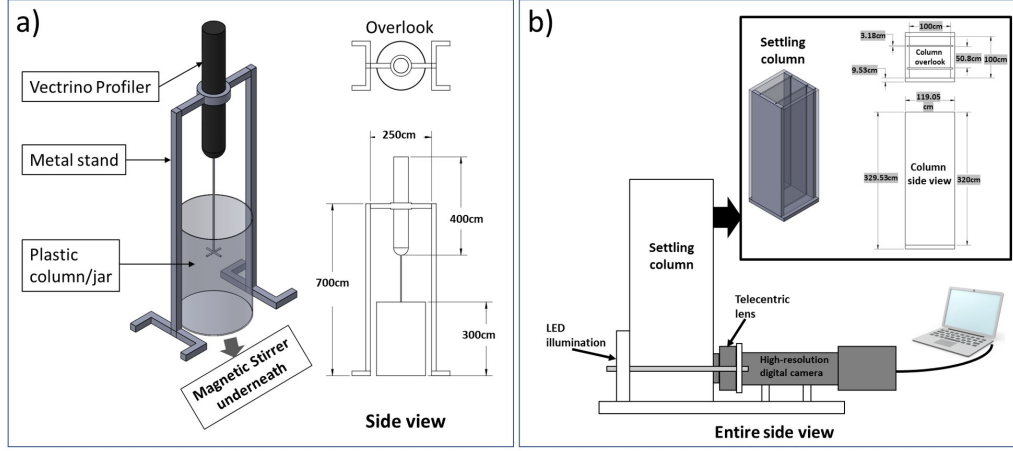


Figure 1: Schematic of the laboratory experimental setup. a) shows the self-designed Vectrino stand set and b) show the LabSFLOC-2 system

Table 1: Various proportions of mineral clay and oil in each experimental run.

Sample	Saline water(L)	Kaolinite clay(g)	Bentonite clay(g)	Texas crude oil(g)
S01	1.00	0.50	/	/
S02	1.00	0.50	/	1.00
S03	1.00	/	0.50	/
S04	1.00	/	0.50	1.00
S05	1.00	0.25	0.25	/
S06	1.00	0.25	0.25	1.00

207 The mass settling velocity of OMAs were observed using the low intrusive
208 LabSFLOC-2 system (the 2nd version of Laboratory Spectral Flocculation
209 Characteristics instrument) (Figure 1b). This instrument was originally de-
210 veloped by Manning and Dyer (2007) and first used by Gratiot and Man-
211 ning (2004). It measures the entire floc population for each sample being
212 assessed and has been successfully applied in many cohesive sediment trans-
213 port studies (Manning et al., 2010; Manning and Schoellhamer, 2013; Uncles
214 and Mitchell, 2017). LabSFLOC-2 utilizes a low-intrusive 2.0 MP Grasshop-
215 per monochrome digital video camera to optically observe individual flocs
216 (e.g. Manning and Dyer (2002)) as they settle in a 350 mm high by 100 mm
217 square Perspex settling column. The video camera, positioned nominally 75
218 mm above the base of the column, views all particles in the center of the
219 column that pass within a 1 mm depth of field, 45 mm from the Sill TZM
220 1560 high-magnification ($5\text{ }\mu\text{m}$ pixel resolution) Telecentric (maximum pixel
221 distortion of 0.6 %), 0.66 (1:1.5) magnification, F4, macro lens fitted behind
222 a 5 mm thick glass faceplate. The LabSFLOC-2 settling column sampling
223 was conducted at the end of each OMA sample’s experiment and these OMA
224 samples are believed in an equilibrium stage.

225 A high-resolution digital microscope system has been used to observe
226 detailed floc structures and to carry out statistical analysis on floc numbers in
227 order to evaluate flocculation rate. All the floc samples were directly collected
228 from the running experiment in real-time using wide mouth ($> 2\text{ mm}$) plastic
229 pipettes to minimize floc disturbance and to transfer the samples from the

230 mixing jar to the microscope slides without using coverslip to prevent the
 231 samples being squeezed. Floc samples are observed with a 4~10 times zoom-
 232 in screen on a DELL laptop by the camera software provided by AmScope
 233 Inc.

234 2.2. Data processing

235 2.2.1. LabSFLOC-2 camera floc data

236 As one of the most commonly used floc video camera instruments in es-
 237 tuarine and coastal suspended sediment transport study, the LabSFLOC-2
 238 produces not only visible floc individual images but also other essential quan-
 239 titative floc properties including floc size, floc shape and floc settling velocity
 240 ([Manning et al., 2010](#)). Through additional theories, other floc quantities can
 241 be derived, such as floc density, fractal dimension and so on. The recorded
 242 videos of floc settling videos can be analyzed with Matlab software routines
 243 based on the HR Wallingford Ltd DigiFloc software ([Benson and Manning,](#)
 244 [2013](#)) and Java Script to semi-automatically process the digital recording
 245 image stack to obtain floc size and settling velocity spectra ([Manning et al.,](#)
 246 [2010](#); [Uncles and Mitchell, 2017](#)). Using the measured floc diameter D (floc
 247 sphere-equivalent diameter), settling velocity W_s , and floc shape, a modified
 248 Stokes Law ([Stokes, 1851](#)) is used to estimate individual floc effective density
 249 ([Manning and Schoellhamer, 2013](#)):

$$\rho_e = \frac{18W_s\beta}{\alpha} \cdot \frac{\rho\nu}{g} \cdot \frac{1 + 0.15Re^{0.687}}{D^2} \quad (1)$$

in which ρ is the saltwater density, ν is the kinematic viscosity, g is gravitational acceleration, α and β are shape-related coefficients and they equal to 1 for perfect spheres. Re in the equation is the particle Reynolds number, defined as

$$Re = W_s D / \nu \quad (2)$$

This modified Stokes' Law is used for encountering flocs which have particle Reynolds numbers greater than its original unity. By assuming floc has a fractal structure, the fractal dimension of floc (n_f) can be calculated via the following relationship (Winterwerp and Van Kesteren, 2004):

$$\left(\frac{D}{d}\right)^{n_f-3} = \frac{\rho_e}{\rho_s - \rho} \quad (3)$$

in which d is the minimum primary particle size which is assumed as ??? (Manning and Schoellhamer (2013)), ρ is the seawater density.

2.2.2. Microscope images analysis

The floc images (e.g., Figure 2a) collected from the digital microscope of each floc sample allow a detailed investigation of floc and OMA structures and the red contours point out the individual flocs which can be manually selected under the screen by real-time observation. Microscope images also provide independent and high-resolution data of floc population, which are number counted manually according to the contours, and shape analyzed for further statistical analysis and flocculation rate evaluation. For each sample,

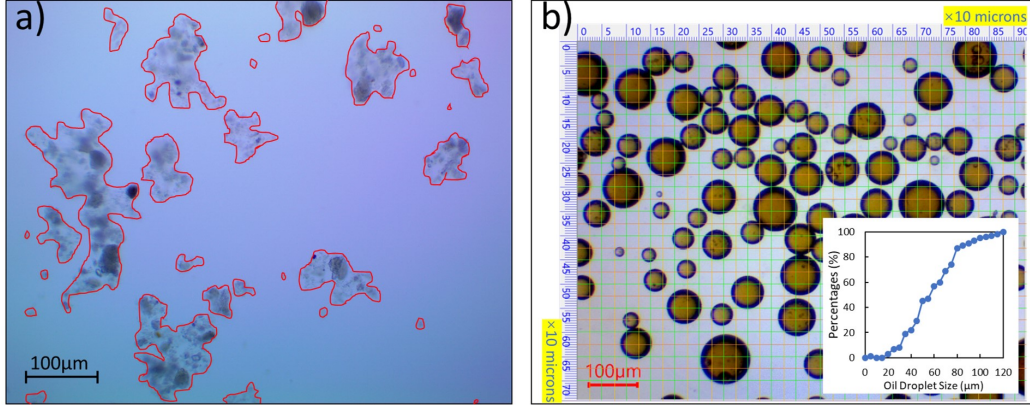


Figure 2: Examples of digital microscope images. a) shows floc number manually counted from a microscope screenshot. The red parts indicate the individual floc can be recognized manually for further statistical analysis. b) shows oil droplets can be manually counted to give the averaged droplet size. The oil droplets are formed and measured after 20 minutes stirring under given turbulence.

268 six different microscopy images have been analyzed which cover hundreds to
 269 thousands individual flocs.

270 We also use microscope images to calculate averaged oil droplets size
 271 under the given turbulence level (e.g., see Figure 2b). The statistical analysis
 272 of the pure oil droplets samples images shows the maximum oil droplets
 273 size can be up to 120 microns and the mean droplets size is approximately
 274 57 microns. Pure oil droplets size distribution under the constant given
 275 turbulence and salinity has been provided in Figure 2b.

276 2.2.3. Mineral stickiness quantification (Flocculation rate (R_f))

277 In this study, we quantify the effect of different mineral clays on OMA
 278 flocculation by their stickiness. In the study of flocculation with significant
 279 organic content, such as due to the presence of transparent exopolymer par-

280 ticles (TEP) (e.g., [Passow \(2002\)](#)), the stickiness can be quantified by per-
 281 forming experiments to estimate flocculation (efficiency) rate [Engel \(2000\)](#).
 282 We perform flocculation rate experiment for three types of mineral particles
 283 without the presence of oil, namely cases S01 (Kaolinite), S03 (Bentonite)
 284 and S05 (half-half mixture of Kaolinite and Bentonite) (see [Table 1](#)), respec-
 285 tively. Temporal microscopy images (six images for each sample at a time)
 286 have been collected during mineral flocs development in a magnetic stirrer
 287 jar from beginning (0 minute) to the end (2 hours) for each mineral sample.
 288 By counting all the floc numbers and normalizing by maximum floc number
 289 of each mineral sample which covers hundreds to thousands individual flocs,
 290 flocculation evolution time series has been obtained for each type of min-
 291 eral clay ([Figure 3](#)). The manually counted floc numbers cover hundreds to
 292 thousands individual flocs which are statistically significant to represent the
 293 entire floc distribution and characteristics of each tested sample.

294 Due to flocculation, the number of particles in each case decays in time. In
 295 the semi-logarithmic plot shown in [Figure 3](#), we observed a nearly exponen-
 296 tial decay of particle number in the first couple of minutes of the flocculation
 297 before the particle number becomes more or less constant in time. By fitting
 298 the first three data points in each run, we obtain the representative floccu-
 299 lation rate: R_f . The three trend lines in [Figure 3](#) indicate that Kaolinite
 300 clay has the lowest flocculation rate of $R_{f_Kaolinite} = 0.13 \text{ (min}^{-1}\text{)}$ while the
 301 Bentonite clay shows a high flocculation rate of $R_{f_Bentonite} = 0.66 \text{ (min}^{-1}\text{)}$
 302 nearly 5 time larger. The mixture of equal amount of Kaolinite and Bentonite

303 has an intermediate flocculation rate of $R_{f_mixed} = 0.32 \text{ (min}^{-1}\text{)}$. Following
304 [Engel \(2000\)](#), we will consider Kaolinite having the lowest cohesion, followed
305 by the mixture of Kaolinite-Bentonite and Bentonite is among the most co-
306 hesive sediments investigated in this study. Noticeably, in the Bentonite (red
307 dots) and mixture (yellow dots) samples, floc numbers increase slightly after
308 reaching the maximum flocculation (lowest normalized floc number) at 5~6
309 minutes and reach to equilibrium stage after 30 minutes, which may imply
310 the break-up of part of the larger fragile flocs. This indicates that the partici-
311 pant of kaolinite not only reduces the cohesion of pure bentonite, meanwhile,
312 the existence of bentonite weakens the stability of pure kaolinite flocs.

313 The primary goal of this mineral flocculation rate quantification present
314 here is to compare the difference of stickiness (not floc number) of bentonite
315 and kaolinite because floc numbers between bentonite and kaolin cases have
316 huge difference (bentonite aggregates much larger but less numbers of flocs
317 because of the much higher stickiness) although bentonite sample and kaolin
318 sample have equivalent mineral concentration under same turbulence condi-
319 tion. This information will be shown later to be very useful for the interpre-
320 tation of the OMA structure and LabSFLOC-2 settling column experimental
321 results.

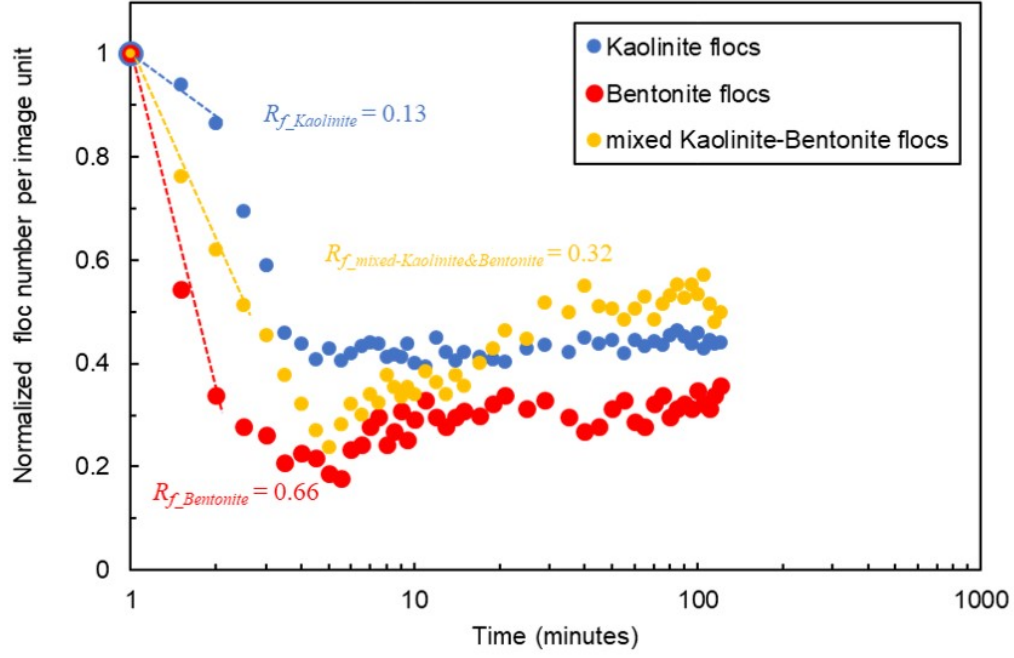


Figure 3: Statistical results of temporal (120 minutes) floc evolutions of normalized floc number for Kaolinite run (blue, S01), Bentonite run (red, S03) and mixed Kaolinite-Bentonite run (yellow, S05). The initial particle number of each run is used for normalization. Each data point comes from manually counted floc number from six different images of each pipette sample which covering hundreds to thousands individual flocs and normalized by the maximum floc number (Kaolinite:1260, Bentonite:782 and Mixture:1323) of each sample.

3. Results

3.1. Floc structures

Samples from each case presented in Table 1 were collected after the flocculation reached equilibrium and have been analyzed by microscopy. High-resolution images provided the details of each mixture floc structure with a magnification factor of 10. Three basic floc types have been observed depending on mineral type: pure mineral flocs/aggregates (no oil) (Figure 4a₁, b₁ & c₁), oil droplets attaching/combining Kaolinite aggregates (Figure 4a₂₋₄), and large flake shaped oil-Bentonite aggregates (Figure 4b₂₋₄).

Figure 4a₁ shows a representative microscope image of settled pure Kaolinite clay flocs (500 mg·l⁻¹, S01 in Table 1). With the addition of 1 g Texas crude oil (S02 in Table 1), the oil droplets can be observed being attached or embraced within the Kaolinite clay structures (such as Figure 4a₂₋₄). The attachment is limited to the surface of oil droplets while the droplets structure remains intact. The oil-Kaolinite aggregates observed are consistent with the droplet OMA type reported in the previous studies (such as Stoffyn-Egli and Lee 2002; Khelifa et al. 2002), in which oil droplets are coated by sediment aggregates through surface attachment. The quantity of mineral attached to a droplet is highly variable.

The OMA obtained from the Bentonite clay run (S03 in Table 1), generated with same turbulent dissipation rate, are shown in Figure 4b₁₋₄. The Bentonite flocs are generally larger than Kaolinite flocs and their size can be up to 100~200 µm in width and several hundred micrometers in length (Figure

345 4b₁). These features are distinct from the pure Kaolinite run (S01) shown in
 346 Figure 4a₁. As demonstrated in section 2.2.3 (or Figure 3), pure Bentonite
 347 clay particles are much cohesive and attachable than Kaolinite particles.
 348 More importantly, the more cohesive characteristic of Bentonite floc leads
 349 to an entirely re-shaped oil-mineral structure (see Figure 4b₂, S04 in Table
 350 1). Compared with oil-Kaolinite flocs, the sphere-shaped oil droplets disap-
 351 peared, and the oil-Bentonite flocs show much larger size of oil soaked mineral
 352 having a flake-shaped aggregates up to hundreds of microns in size (see Fig-
 353 ure 4b₂₋₄). Compared with the previous studies (Stoffyn-Egli and Lee, 2002;
 354 Khelifa et al., 2002; Zhao et al., 2016), the dominant oil-Bentonite aggre-
 355 gates observed here belong to a dense type of oil-aggregate called flake/solid
 356 OMA. Flake aggregates have the appearance of membrane structures, usu-
 357 ally floating or neutrally buoyant, which can attain hundreds of microns in
 358 length. Their microstructure is highly organized as dendritic or feather-like.
 359 Experimental results suggest that high shear strength (i.e. extended or faster
 360 agitation) tends to break or crumple flake aggregates. The crumpled flakes
 361 (Figure 4b₂₋₄) may be distinguished from mineral-embraced droplet OMA
 362 (Figure 4a₂₋₄) by their folds or preferential orientation of the minerals.

363 After mixing equal amount of Kaolinite and Bentonite clay for Case S05,
 364 the mixture flocs contain both Kaolinite floc and Bentonite floc structures
 365 (see Figure 4c₁), and importantly, although the general size of the mixed
 366 flocs (Figure 4c₁) are smaller than pure Bentonite case (Figure 4b₁). The
 367 Bentonite floc structure appears to be dominant in the mixture mineral sam-

ple. With the addition of oil in the mixture sample S06, large flake shaped OMA can be observed in Figure 4f which has similar floc size with those in the oil-Bentonite case (S04, see Figure 4b₂₋₄). However, both droplet OMAs and flake OMAs can be observed (Figure 4c₃₋₄).

Due to different cohesion between Kaolinite and Bentonite clays, the resulting OMA structures are also distinctly different which is expected to lead to different settling velocities. In the next sections, we will investigate different mineral flocs and OMAs settling velocity and discuss their relationship to floc structures.

3.2. Floc physical properties

The previous section provided insights on the floc structures for different types of OMAs by microscopy images. This section is devoted to more quantitative study of floc physical properties, particularly their settling velocities. The scatterplots in Figure 5a, 6a & 7a illustrate individual spherical-equivalent dry mass weighted floc sizes (x-axis) plotted against their corresponding settling velocities (y-axis) of each sample (see Table 1) collected and analyzed by LabSFLOC-2 camera system. The scatterplots allow subsequent statistical analysis for floc properties using 12 different size classes (Size Band details are shown in the bottom of Figure 5-7). The physical properties of particular interest here are, the counted floc numbers of each size band (Figure 5c, 6c & 7c), the settling velocity (Figure 5d, 6d & 7d, floc density (Figure 5e, 6e & 7e) and fractal dimension (Figure 5f, 6f & 7f).

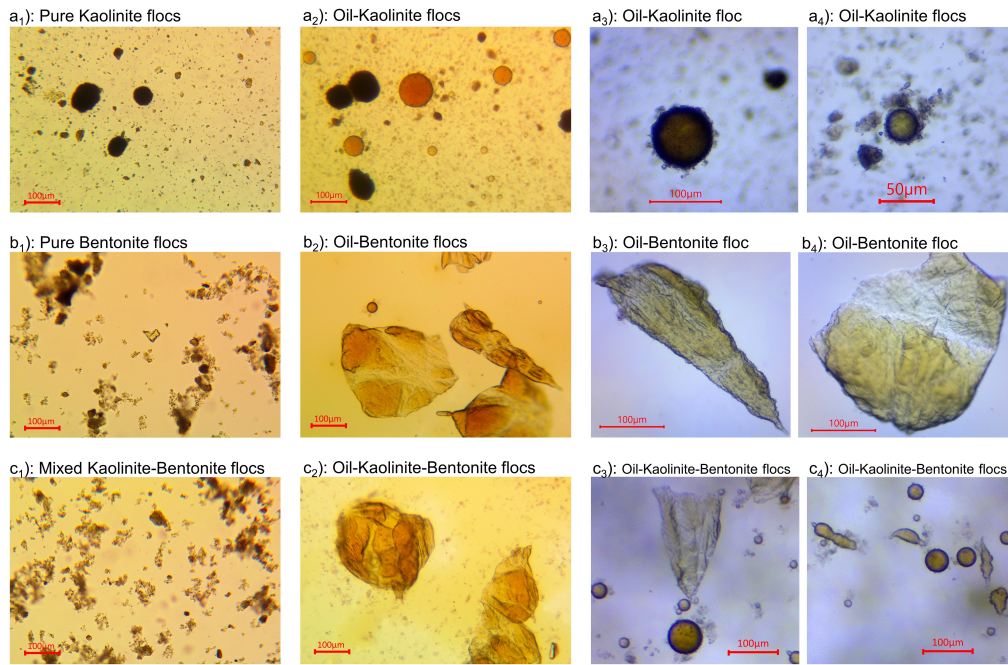


Figure 4: Floc images from the high-resolution digital microscope camera. a1-a4) Kaolinite (S01) and oil-Kaolinite (S02) samples; b1-4) Bentonite (S03) and oil-Bentonite (S04) samples; c1-4) mixed Kaolinite and Bentonite (S05) and oil-Kaolinite-Bentonite (S06) samples.

3.2.1. Kaolinite and Oil-Kaolinite flocs

The scatterplot presented in Figure 5a indicates that the Kaolinite clay flocs (S01) cover a size range from 20 to 400 microns while their settling velocities vary from 0.04 to 10 mm·s⁻¹. Although with some notable scatters, the floc settling velocities are more or less proportional to floc sizes. Adding oil into Kaolinite mineral (S02) shows negligible change in the floc size range (see Figure 5b). However, when floc size is greater than about 80 micron, the peak settling velocities (about 4 mm·s⁻¹) show almost no change with respect to floc size. Also, there exist some low-density flocs in the rather large size range of 200~400 microns with settling velocities ranging from 0.2~0.6 mm·s⁻¹ (around the red constant density line of 16 kg·m⁻³ in Figure 5b). This is due to the large Oil-Kaolinite flocs having much lower density than those of pure Kaolinite flocs. A more quantitative understanding on these interesting features can be obtained by examining the statistics of 12 size bands.

The number of Kaolinite flocs increases dramatically from Size Band (SB)-1 (20~40 microns) to SB-3 (80~120 microns) and then drops quickly from SB-3 to SB-8 (320~400 microns) (Figure 5c, blue bands). Adding oil to Kaolinite significantly increases floc number for small size flocs (20~80 microns) at SB-1 and SB-2 while floc numbers at larger size class are generally lower than or similar to those of pure Kaolinite flocs (Figure 5c, orange bars). The settling velocities of Kaolinite samples (S01 and S02 in Table 1) averaged for each size class are shown in Figure 5d. Evidently, pure mineral

flocs (S01) show a rapid increase of settling velocities with respect to the increase of floc sizes for the entire size class spectrum (SB-1 to SB-8). On the contrary, oil-Kaolinite flocs show milder increase of settling velocity with respect to floc size from SB-1 to SB-6 until a completely different trend is observed for larger size class (SB-6 to SB-8), namely, a significant decrease of settling velocity with respect to increase of floc size. Overall, adding oil to Kaolinite decreases flocs settling velocity, particularly for larger size classes (by nearly factor 3 in the SB-6 and nearly a factor 7 in the SB-8). Considerable reduction of settling velocity at SB-6 to SB-8 is clearly associated with the significant decrease of floc effective density due to the addition of oil to Kaolinite at this size range (see Figure 5e). Generally, adding oil reduces floc effective density in all size ranges of flocs but the reduction is much more pronounced at large size class. In SB-2 and SB-3, effective density decreases by approximate $1/3$ to $1/4$ by adding oil while settling velocity also decreases by $1/3$ to $1/4$. In SB-4 to SB-8, the effective density decreased by half or much more especially in the largest sized flocs, and their settling velocity shows a remarkable reduction in large size flocs such as SB-6 to 8. Since averaged droplet size is about 57 microns as measured in the laboratory tests, it is very likely that there is less oil contained in smaller flocs size classes. Overall, the results presented here is consistent with the presence of oil as droplets (see Figure 4a₂-a₄) having lower density than saltwater or mineral. The low density oil droplets contribute to the reduced settling velocity (or floc density) particularly at the large floc size range.

436 The fractal dimension for Kaolinite flocs or Kaolinite-oil flocs is in the
437 range of 2.4~2.6 except for a small number of large flocs in SB-7/8. In gen-
438 eral, adding oil slightly reduces fractal dimension to 2.4. A notable exception
439 is that when oil is added to Kaolinite, the largest flocs in SB-8 show a much
440 lower fractal dimension of 2.05 due to containing low density oil droplets in
441 the large structure.

442 3.2.2. Bentonite and Oil-Bentonite flocs

443 In the pure Bentonite sample (S03, Figure 6a), we observe some very
444 large size flocs up to 400~700 microns that do not exist in the pure Kaolinite
445 sample (S01). The resulting settling velocity range is also wider (0.01~20
446 mm·s⁻¹) than that in Kaolinite samples. A more careful examination further
447 suggests that many large size flocs (in SB-9 SB-12 in Figure 6c, 400~700
448 microns) in pure Bentonite sample (S03) are of very low density (within 50
449 kg·m⁻³) and their settling velocities are limited to range of 1~5 mm·s⁻¹,
450 despite very large floc size. Importantly, adding oil further increased the
451 floc size up to 800 microns (Figure 6b), and it also shows an upper limit of
452 settling velocity but at much higher value of about 10 mm·s⁻¹ compared to
453 that of oil-Kaolinite (see Figure 5b).

454 Quantitatively, floc number increases from SB-1 to SB-5 and then reduces
455 afterwards to SB-12 (floc number < 5 in SB-12) (Figure 6c). Comparing to
456 Kaolinite samples (S01 and S02), the most notable difference is that the floc
457 numbers for Bentonite samples are significantly lower than those in Kaoli-

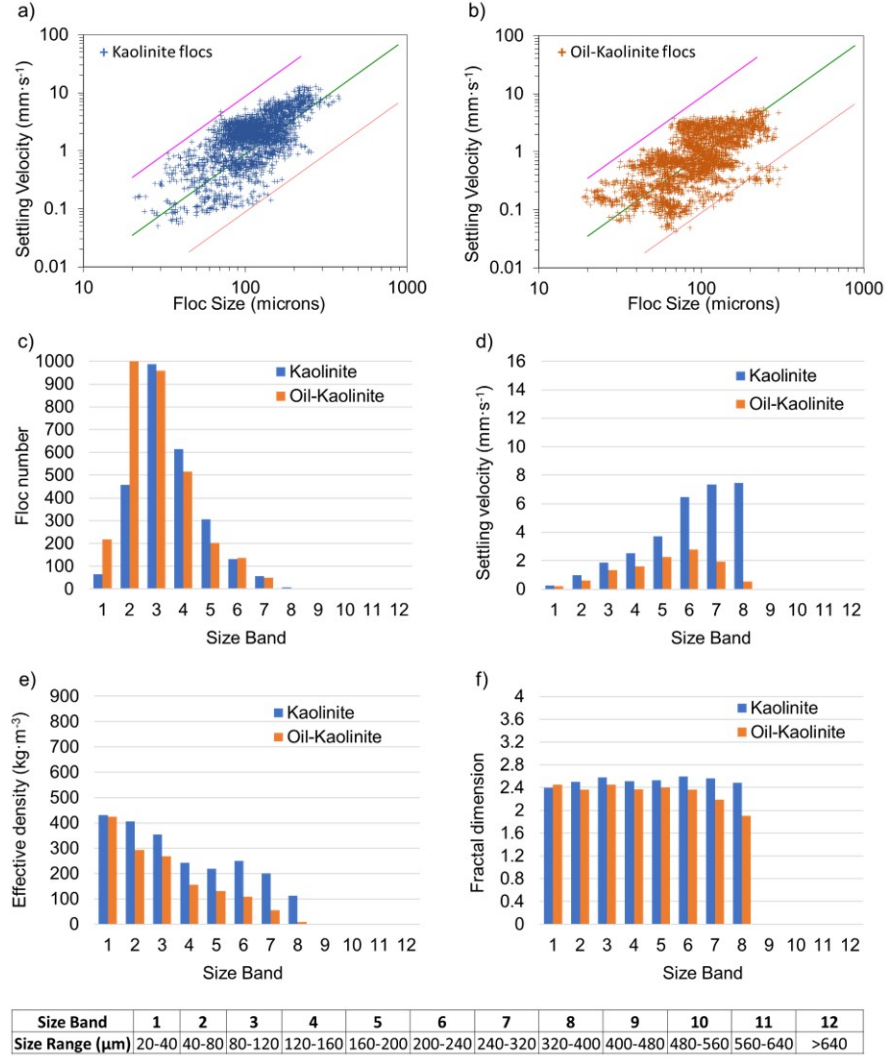


Figure 5: Floc characteristics from the LabSFLOC-2 analysis. a) & b) show the plots of floc sizes vs. settling velocities of Kaolinite (S01) and Oil-Kaolinite (S02), respectively. The three diagonal lines present contours of Stokes settling velocity calculated with a constant effective density (i.e. floc bulk density minus water density) of $1,600 \text{ kg}\cdot\text{m}^{-3}$ (pink line, equivalent to a quartz particle), $160 \text{ kg}\cdot\text{m}^{-3}$ (green) and $16 \text{ kg}\cdot\text{m}^{-3}$ (red line). c)-f) show the 12 Size Bands trends of floc number, settling velocity, floc density and fractal dimensions respectively for both Kaolinite (blue) and Oil-Kaolinite (orange) samples.

nite sample. Larger floc sizes and lower floc number in Bentonite samples
 are consistent with the high flocculation rate (high stickiness) of Bentonite
 discussed in Section 2.2.3. When oil is added to Bentonite (S04) we observe
 an increase of settling velocities with the floc size (Figure 6d), except at the
 largest size class (SB-12). This trend is more or less consistent with pure
 Kaolinite sample (S01, Figure 5d) and Bentonite sample (S03, Figure 6d),
 but different from the oil-Kaolinite sample (S02, Figure 5d). This suggests
 that oil interact differently with Kaolinite and Bentonite samples and it is
 consistent with their distinct droplet OMA and flake OMA structures pre-
 sented in Figure 4. More importantly, when oil is added to Bentonite, we
 observe a more rapid increase of settling velocity when floc size increases from
 SB-9 to SB-12. Furthermore, comparing to the pure Bentonite condition, we
 obtain an increase of settling velocity by more than a fact of 2 in SB-11, while
 recall that for Kaolinite samples, adding oil to Kaolinite (S02) significantly
 reduces the floc settling velocity. These observations are supported by the
 data from floc effective density. From Figure 6e, we can see that adding oil
 to Bentonite clay generally increases floc effective density with the most sig-
 nificant increases occur at SB-1 and SB-9 to SB-12 (contrast with Figure 5e,
 adding oil reduces floc effective density in Kaolinite samples). In this case,
 the oil droplets no longer exists and become absorbed into mineral flocs. It
 is likely that at such micro-scale, oil changes the adhesion characteristic and
 make the small flocs more compact, dense with lower porosity.

The fractal dimension for Bentonite floc or Bentonite-oil flocs are in the

range of 2.2~2.4 which is slightly lower than those of Kaolinite samples. However, adding oil to Bentonite generally increases fractal dimension with the largest increase occurs at SB-1 with a fractal dimension near 2.5. It is interesting to the point out that, a notable fractal dimension changes after adding oil is in larger size class floc of SB-8 (320~400 microns) for Kaolinite sample and in the smallest size class of SB-1 (20~40 microns) for Bentonite sample. This drastic difference is again consistent with different OMA structure of Kaolinite and Bentonite clays.

3.2.3. *Mixed Kaolinite-Bentonite and Oil-Kaolinite-Bentonite flocs*

In the mixed Kaolinite-Bentonite sample (Figure 7a), a large number of small sized flocs (<80 microns) are observed which are lacking in pure Kaolinite (Figure 5a) or pure Bentonite (Figure 6a) samples. The significant portion of small flocs in SB-12 may due to in the mixture, the kaolinite particles decrease the stickiness of the mixed floc comparing to pure bentonite condition, which may also lead to more small bentonite flocs. Meanwhile, it is likely that the bentonite flocs, when flocculate with kaolinite flocs, make the whole mixed flocs more fragile and with the high turbulence level provided, part of the mixture flocs tends to break-up into smaller flocs, even smaller than those in pure bentonite or kaolinite cases. The observed large amount of microflocs in Figure 7 at the equilibrium stage is consistent with the temporal evolution of normalized floc number shown in Figure 3 that at the later stage, normalized floc number increases. Therefore, when mixing low stickiness

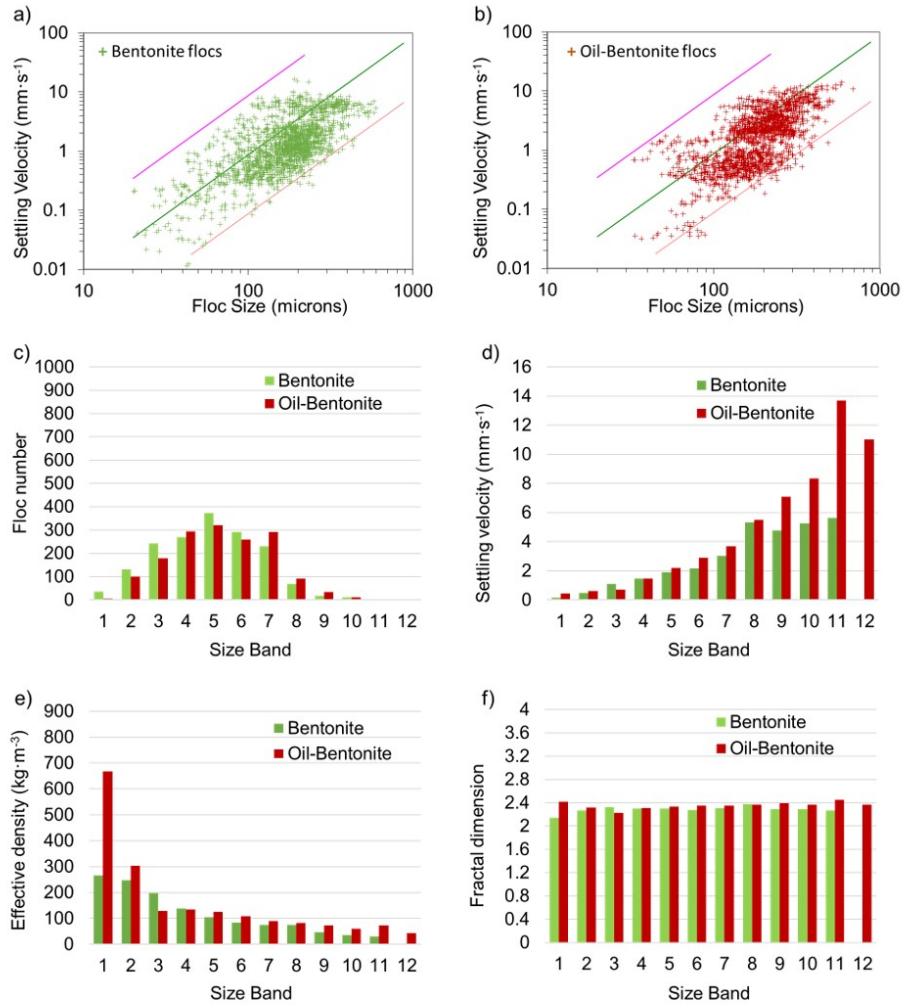


Figure 6: Floc characteristics from the LabSFLOC-2 analysis. a) & b) show the plots of floc sizes vs. settling velocities of Bentonite (S03) and Oil-Bentonite (S04), respectively. The three diagonal lines present contours of Stokes settling velocity calculated with a constant effective density (i.e. floc bulk density minus water density) of $1,600 \text{ kg}\cdot\text{m}^{-3}$ (pink line, equivalent to a quartz particle), $160 \text{ kg}\cdot\text{m}^{-3}$ (green) and $16 \text{ kg}\cdot\text{m}^{-3}$ (red line). c)-f) show the 12 Size Bands trends of floc number, settling velocity, floc density and fractal dimensions respectively for both Bentonite (blue) and Oil-Bentonite (orange) samples.

503 kaolinite and high stickiness bentonite, the flocculation process is much more
504 complex and it takes longer time to reach equilibrium. This is because when
505 more porous/fragile bentonite floc structure initially combined with denser
506 kaolinite floc, the entire mixture flocs become fragile to high turbulence and
507 breaks into more small flocs.

508 The settling velocity for pure Kaolinite-Bentonite flocs peaks at about
509 $10 \text{ mm}\cdot\text{s}^{-1}$ for floc size greater than about 100 microns. When oil is further
510 added to the mixed Kaolinite-Bentonite sample (Figure 7b), we observe even
511 higher settling velocity flocs exceeding $10 \text{ mm}\cdot\text{s}^{-1}$, which is clearly due to
512 higher floc effective density in SB-3 to SB-9 shown in Figure 7e. We also
513 obtain more small flocs but their density is widely spread from nearly close
514 to water (below the red line of $16 \text{ kg}\cdot\text{m}^{-3}$) to those high-density flocs (between
515 the green line of $160 \text{ kg}\cdot\text{m}^{-3}$ and red line of $1600 \text{ kg}\cdot\text{m}^{-3}$).

516 Generally, both Kaolinite-Bentonite minerals flocs and Oil-Kaolinite-Bentonite
517 flocs show increasing settling velocity with the increasing floc sizes (Figure
518 7d) except at the largest size class. In other words, by adding oil to equally
519 mixed Kaolinite and Bentonite mixture, the overall settling velocity trend is
520 similar to that of pure Bentonite (Figure 6d). This observation can be further
521 confirmed by examining floc effective density shown in Figure 7e. Similar to
522 adding oil to pure Bentonite (see Figure 6e), adding oil to Kaolinite-Bentonite
523 mixture generally increase floc effective density and hence the settling veloc-
524 ity also increases in most Size Bands. A minor difference is that the rise
525 of settling velocity in SB-9 to SB-12 by adding oil is less dramatic in SB-9

526 to SB-12 in the Kaolinite-Bentonite mixture case. Also, there is simply less
527 number of those large flocs in Kaolinite-Bentonite mixture cases, suggesting
528 that adding oil to Kaolinite-Bentonite mixture does not increase the cohe-
529 sion as much when comparing to adding oil to pure Bentonite, possibly due
530 to the presence of less cohesive Kaolinite. Therefore, oil can be considered
531 to preferably interacting with Bentonite and the presence of Kaolinite is of
532 secondary effect to slightly reduce cohesion.

533 For Kaolinite-Bentonite mixture, the fractal dimension of small sized flocs
534 can be up to 2.8 while the large flocs are of lower value around 2.4 to 2.6.
535 The range of fractal dimension is larger than pure Bentonite (S03) and it is
536 similar to pure Kaolinite (S01) except for the smallest size class (SB-1). By
537 adding oil, flocs fractal dimension in larger size class SB-6 to SB-9 increases
538 and those in small size class SB-1 to SB-5 show negligible change.

539 *3.3. Microflocs and Macroflocs*

540 In the cohesive sediment literature, two distinguished floc components:
541 microflocs and Macroflocs, have been utilized to quantitatively describe the
542 floc spectra ([Manning et al., 2010](#); [Manning and Dyer, 2007](#); [Manning and](#)
543 [Schoellhamer, 2013](#)), and a floc diameter of 160 μm has been often used to
544 distinguish between microflocs and Macroflocs groups ([Manning and Dyer,](#)
545 [2002](#); [Manning, 2004](#); [Manning et al., 2010](#)). In order to obtain more gen-
546 eral understanding on the floc physical properties, a summary of mean floc
547 properties for the entire floc population and sub-population categorized into

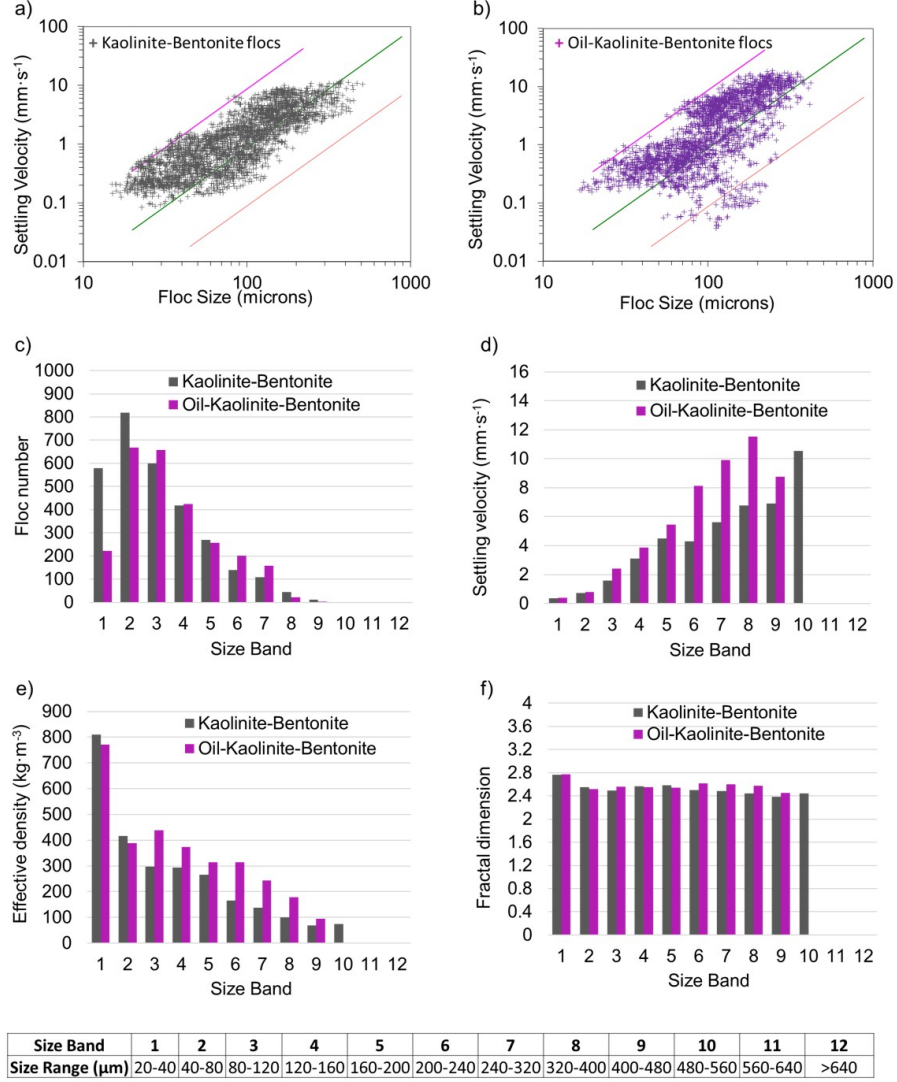


Figure 7: Floc characteristics from the LabSFLOC-2 analysis. a) & b) show the plots of floc sizes vs. settling velocities of mixed Kaolinite-Bentonite (S05) and Oil-Kaolinite-Bentonite (S06), respectively. The three diagonal lines present contours of Stokes settling velocity calculated with a constant effective density (i.e. floc bulk density minus water density) of $1,600 \text{ kg}\cdot\text{m}^{-3}$ (pink line, equivalent to a quartz particle), $160 \text{ kg}\cdot\text{m}^{-3}$ (green) and $16 \text{ kg}\cdot\text{m}^{-3}$ (red line). c)-f) show the 12 Size Bands trends of floc number, settling velocity, floc density and fractal dimensions respectively for both mixed Kaolinite-Bentonite (blue) and Oil-Kaolinite-Bentonite (orange) samples.

Table 2: Summary of microfloc and Macrofloc mean quantities of each sample investigated in this study.

Samples	Kaolinite			Bentonite			Kaolinite-Bentonite		
Demarcation	Total	micro	Macro	Total	micro	Macro	Total	micro	Macro
N	2631	2128	503	1705	681	1024	2998	2420	578
$\bar{D}(\mu\text{m})$	120	101	199	185	105	238	104	76	225
$\bar{\rho}_e(\text{kg}\cdot\text{m}^{-3})$	315	336	224	127	187	87	410	461	200
$\bar{W}_s(\text{mm}\cdot\text{s}^{-1})$	2.41	1.82	4.89	2.00	1.08	2.61	1.97	1.27	4.90
\bar{f}_n	2.54	2.54	2.55	2.30	2.30	2.30	2.57	2.59	2.53

Samples	Oil-Kaolinite			Oil-Bentonite			Oil-Kaolinite-Bentonite		
Demarcation	Total	micro	Macro	Total	micro	Macro	Total	micro	Macro
N	3102	2696	406	1592	580	1012	2610	1975	635
$\bar{D}(\mu\text{m})$	102	86	204	198	115	246	120	87	222
$\bar{\rho}_e(\text{kg}\cdot\text{m}^{-3})$	249	269	113	127	167	104	408	446	290
$\bar{W}_s(\text{mm}\cdot\text{s}^{-1})$	1.21	1.03	2.41	2.53	1.07	3.36	3.33	1.94	7.63
\bar{f}_n	2.40	2.40	2.36	2.33	2.29	2.35	2.57	2.57	2.58

microflocs and Macroflocs for all cases are presented in Table 2. Similar to the previous section, the physical floc properties of interest here are floc number (N), mean floc size (\bar{D}), mean effective density ($\bar{\rho}_e$), mean settling velocity (\bar{W}_s) and mean fractal dimensions (\bar{f}_n).

3.3.1. Mineral types influence on flocculation

A comparison between Kaolinite (S01) and Bentonite flocs (S03) show that the Kaolinite flocs in total have around 35 % higher N (2631 versus 1705) and 35 % smaller \bar{D} (120 versus 185 μm) (see Table 2). Moreover, the larger N in Kaolinite is only due to microflocs and Kaolinite has only half of

557 that of Bentonite in terms of Macrofloc numbers. This matches with the pre-
 558 vious clay flocculation studies in saline water (e.g., [Zhang et al. 2019](#)). Due
 559 to the water salinity, the attraction forces (London-van der Waals' forces)
 560 between the clay plates can be dramatically enhanced over repulsion forces.
 561 Particularly for Bentonite clay, the enhancement of flocculation can be sig-
 562 nificantly larger than Kaolinite does under the same salinity which results
 563 in more Macroflocs forming in Bentonite case. The $\bar{\rho}_e$ for the entire floc
 564 population of Kaolinite floc is around 2.5 times higher than that of Ben-
 565 tonite floc. This is particularly due to the Macroflocs with significantly low
 566 $\bar{\rho}_e$ (only $87 \text{ kg}\cdot\text{m}^{-3}$) in Bentonite with a mean density of only $87 \text{ kg}\cdot\text{m}^{-3}$.
 567 Despite somewhat smaller \bar{D} of Kaolinite flocs, their significantly larger $\bar{\rho}_e$
 568 results in approximately 20 % larger \bar{W}_s than that of Bentonite flocs. Finally,
 569 \bar{f}_n of Kaolinite flocs is about 2.54, which is higher than that of Bentonite
 570 flocs of around 2.3. The differences of Kaolinite and Bentonite flocs revealed
 571 here can be directly link to the different types of mineral particle stickiness
 572 in saline water (see Figure 3) which shows Bentonite have almost five times
 573 higher R_f than that of Kaolinite in the seawater of 35 ppt. Even with the
 574 much lower salinity condition (1.2~3.5 ppt, tested by [Khelifa et al. \(2005b\)](#)),
 575 the clay types still can be a significant factor for influencing the flocculation.

576 The microfloc \bar{D} in the mixed sample is 15 % smaller than that of Kaoli-
 577 nite, and yet in terms of the Macrofloc \bar{D} , mixed sample is 13 % larger.
 578 In fact, the Macrofloc \bar{D} for mixed Kaolinite-Bentonite sample is nearly
 579 comparable (only 5 % smaller) to that of pure Bentonite. Although the

580 resulting \overline{W}_s of mixed Kaolinite-Bentonite sample for the whole floc popu-
 581 lation ($1.97 \text{ mm}\cdot\text{s}^{-1}$) is very similar to that of pure Bentonite sample (2.0
 582 $\text{mm}\cdot\text{s}^{-1}$), the mixed sample reaches the similar \overline{W}_s due to having the largest
 583 $\overline{\rho}_e$ ($410 \text{ kg}\cdot\text{m}^{-3}$) and the smallest \bar{D} (104 microns) comparing with those of
 584 the two pure clay samples. Looking more into the difference, we can see
 585 that Macrofloc \bar{D} and Macrofloc $\overline{\rho}_e$ of Kaolinite-Bentonite mixture is about
 586 13 % larger and 12 % smaller than those of pure Kaolinite sample, respec-
 587 tively, which suggests a slight increase of cohesion in Kaolinite-Bentonite
 588 flocs, possibly due to the presence of Bentonite. In terms of microflocs,
 589 Kaolinite-Bentonite sample show the smallest and the densest flocs which is
 590 also more similar to the microfloc of pure Kaolinite sample, but distinctly
 591 different from those of pure Bentonite sample. Moreover, examining the \bar{D}
 592 of mixed Kaolinite-Bentonite sample further reveals distinct behavior be-
 593 tween microflocs and Macroflocs. The mixture sample possesses a dual fea-
 594 ture, namely, the Kaolinite behavior in microflocs and Bentonite behavior
 595 in Macroflocs. When mixing two types mineral, the Macroflocs development
 596 is slightly enhanced by the more cohesive Bentonite component, but for the
 597 entire mixed floc characteristics the effect of Bentonite appears to be benign
 598 while the effect of Kaolinite appears to be dominant, especially in microfloc
 599 population. This seemingly subtle point between Bentonite and Kaolinite
 600 is raised here because it may play a more important role when interacting
 601 with oil droplets. Particularly for the oil spilled occurring in natural mixture
 602 sediment environment, each clay type influence should be fully understood

603 necessarily.

604 3.3.2. Oil participation in mineral flocculation

605 By adding oil component to the Kaolinite floc sample (S02 in Table 1),
606 oil-Kaolinite N for the entire population increases by around 18 % while the
607 corresponding \bar{D} decreases by around 18 %. However, instead of obtaining
608 a slight increase of $\bar{\rho}_e$ commonly occurs due to decreased cohesion, the $\bar{\rho}_e$
609 for the entire population also decreases by 20 %. As a result of both re-
610 duced \bar{D} and $\bar{\rho}_e$, we obtain a significant reduction of \bar{W}_s by 50 % (decreases
611 from 2.41 mm·s⁻¹ to 1.21 mm·s⁻¹, see Table2). Looking further into the
612 microfloc and Macrofloc statistics, we observe different response of microfloc
613 and Macrofloc due to the addition of oil to Kaolinite in the saline water. The
614 microfloc population shows a 27 % increase in N and 15 % reduction of \bar{D} ,
615 while the Macrofloc population show 20 % reduced in N and very slight 2.5 %
616 increase (or nearly unchanged) \bar{D} . This indicates a small shift to microflocs
617 and reduction of cohesion (flocculation rate) due to the addition of oil. The
618 common and more significant trend for both microfloc and Macrofloc is their
619 reduction of $\bar{\rho}_e$: the microflocs show slight (20 %) decrease of $\bar{\rho}_e$ while the
620 Macroflocs show nearly a factor 2 decrease of $\bar{\rho}_e$. As a result, the microfloc
621 and Macrofloc \bar{W}_s are decreased by 43 % and 51 %, respectively. Overall,
622 the participation of lower density oil droplets reduces the OMA density, con-
623 sistent with the droplet OMA structure presented in Figure 4. The settling
624 velocity data in Table 2 confirms that pure Kaolinite flocs tend to attach

625 with the oil droplets (around the surface) forming OMAs with much lower
626 density than the original pure Kaolinite mineral flocs.

627 For Bentonite clay, adding oil component decreases total N by around 7
628 % while \bar{D} for the entire floc population increases by around 7 %. On the
629 other hand, although the $\bar{\rho}_e$ for the entire floc population is unchanged by
630 adding oil, we obtain 11 % decrease of microfloc $\bar{\rho}_e$ while the Macrofloc $\bar{\rho}_e$
631 is increased more significantly by 20 %. As a result, microfloc \bar{W}_s is nearly
632 unchanged while the Macrofloc \bar{W}_s is increased by 29 %. Since the total flocs
633 are dominated by Macroflocs in the Bentonite cases, the \bar{W}_s for the entire
634 population is increased by 25 % when adding oil mainly and this is caused
635 by the increase of $\bar{\rho}_e$ in Macroflocs.

636 By adding oil into mixed Kaolinite-Bentonite clay (S06), the total N
637 shows a reduction of around 13 %, and total \bar{D} increases by around 15 %.
638 Adding oil decreases microfloc N by 18 % and increases microfloc \bar{D} by 14
639 %. Oil also causes the Macrofloc N to increase by 10 % but with negligible
640 decrease of \bar{D} . Consistent with adding oil to pure Bentonite, here we see a
641 45 % increase in $\bar{\rho}_e$ of Macrofloc by adding oil to mixed Kaolinite-Bentonite
642 sample. As a result, \bar{W}_s increases almost 70 % due to the minor increase of
643 microfloc \bar{D} as well as more dramatic increase of $\bar{\rho}_e$ in Macroflocs.

644 4. Discussion

645 The results presented above indicate unique differences of flocculation
646 characteristics between Kaolinite and Bentonite OMAs which may be closely

647 influenced by the OMA structures. Previously, this hasn't been clearly re-
 648 vealed by studying the natural mixture sediment samples directly (e.g., [Khe-
 649 lifa et al. 2005b, 2008; Sun et al. 2010, 2013; O'Laughlin et al. 2017](#)). Kaoli-
 650 nite particles tend to show lower cohesion and the resulting oil-Kaolinite
 651 aggregates can be categorized as droplet OMAs. The pickering emulsions
 652 ([Chevalier and Bolzinger, 2013](#)) may apply to the oil-Kaolinite droplets flocs
 653 in terms of the single droplet OMA structures (e.g., Figure 4a₂). But the
 654 Kaolinite OMA also can be a bit more complex when multiple mineral flocs
 655 and oil droplets can also possibly attach together and combine to larger
 656 aggregates. In this case, the mineral clay particles/flocs adhere on the oil
 657 droplets surface, and the mineral particles act as a web-structures surround-
 658 ing the oil droplet preventing its attachment to other oil droplets or fur-
 659 ther re-bonding to oil slicks. Previous studies (e.g., [Zhao et al. 2017](#)) have
 660 found that equilibrium droplet oil-sediment aggregates can be considered as
 661 very stable structure and hardly breakup. Since the Kaolinite mineral parti-
 662 cles can be attached together as a much larger structure than individual oil
 663 droplet, the oil can be observed being attached or even embraced within the
 664 Kaolinite flocs (Figure 4a₂₋₄). Because oil droplets structure remains intact,
 665 the oil-Kaolinite OMAs show significantly lower effective density and settling
 666 velocity than the pure Kaolinite flocs. This also apply to the natural clay
 667 mineral formed mixture oil sediment aggregates according to [Khelifa et al.
 668 \(2008\)](#) suggesting that oil-sediment aggregates have 2 to 3 times lower effec-
 669 tive density than pure sediment flocs because of the low density oil droplets

670 attaching within the flocs. On the other hand, Bentonite particles are of
671 very high cohesion and are observed to form large, fluffy (low density and
672 high porosity) and complex aggregate structure. The Bentonite aggregates
673 tend to re-shape and absorb or be absorbed by the oil droplets forming large
674 (as large as 900 microns floc has been observed) dense oil-Bentonite aggre-
675 gates (Figure 4b₂₋₄). Previous experimental work also indicates that the
676 oil-sediment aggregates formed using natural sediment can be as large as
677 900 microns (O’Laughlin et al., 2017) and their settling velocity can be vari-
678 able depending on oil amount trapped in sediment mixture aggregates (Sun
679 et al., 2010). They also suggest that higher sediment concentration may
680 lead to larger flocs with higher settling velocity. In the results presented
681 here further specify the clay type influence to the oil-sediment aggregates’
682 characteristics. Particularly for the Bentonite clay, because the oil droplets
683 no longer exists by themselves and the oil is mainly absorbed at micro-scale
684 level onto the mineral structure, the Bentonite particles can actually become
685 more compact together than its pure mineral floc structure (high porosity)
686 and the resulting OMAs are dominated by denser and larger Macroflocs. The
687 overall settling velocities of oil-Bentonite OMAs are also slightly larger than
688 the pure Bentonite flocs. The response of Kaolinite and Bentonite to the
689 addition of oil are distinctly different due to the corresponding droplet OMA
690 and flake/solid OMA structures, respectively. Overall, in Kaolinite, oil sig-
691 nificantly decreases effective density and hence the settling velocity decreases
692 significantly for the entire floc population. Meanwhile, oil slightly increases

693 settling velocity of oil-Bentonite flocs due to increasing floc size and effective
 694 density in Macrofloc population. Therefore, adding oil into pure Kaolin-
 695 ite has negative influence to the cohesion and flocculation which agrees with
 696 [Khelifa et al. \(2008\)](#). However, the cohesion of Bentonite can be increased by
 697 oil participation because of its unique high porous structure. The stickiness
 698 (cohesion) can be one of the most significant key factor for the OMA floc-
 699 culation. Since chemical dispersant may increase the oil droplets stickiness
 700 to enhance the oil-sediment flocculation after oil spill occurrence in natural
 701 environment ([Khelifa et al. 2008](#)), the higher stickiness of bentonite clay also
 702 may have the similar effect on OMA flocculation positively. The significant
 703 of stickiness influence may also change the relationship between turbulence
 704 level and floc size. Noticeably, both previous study (e.g., [Sun et al. \(2010,](#)
 705 [2013\)](#)) and present study show that in very high turbulence level, turbulent
 706 Kolmogorov length scale has limited effect to the oil-mineral floc size.

707 Furthermore, by adding oil to mixture sample, oil selectively interacts
 708 more actively with Bentonite in Macroflocs rather than with Kaolinite under
 709 a condition of same amount of clay, respectively. The increase of total floc
 710 size and decrease in total floc number by adding oil in the mixed Kaolinite-
 711 Bentonite sample (Table 2) indicate a slight enhance of floc cohesion due to
 712 the addition of oil. However, a more careful observation suggests that the
 713 changes due to the addition of oil are completely different between microfloc
 714 and Macrofloc populations especially in the more significant changes of floc
 715 effective density which directly leads to the settling dynamics difference.

Overall, adding oil transforms the Macroflocs into higher density OMAs with higher settling velocity due to high stickiness Bentonite component. Moreover, oil also increases the density and settling velocity in microflocs dominant by Kaolinite component with less oil participation. Therefore, with the multiple mixed mineral types exist in the natural environments commonly, their own flocculation capability and settling behaviors may vastly different especially when meeting with other contaminant materials such as spilled oil because the third participation during the mineral flocculation processes may totally change the structures of flocs which may directly influence the flocs characteristics such as porosity, density and most importantly settling velocities.

5. Conclusions

To conclude, the LabSFLOC-2 system has been utilized in OMAs studies to understand flocculation characteristics and settling velocities. Droplet OMAs and flake/solid OMAs have been observed in OMA generation which matches with the previous studies. Furthermore, multiple OMAs structures have been studied. For the mixed oil-Kaolin-Bentonite case which is closer to the natural sediment mixture condition, both Kaolin and Bentonite component can be aggregated with oil droplets and develop OMAs to settling. However, Bentonite becomes more dominate in OMA flocculation efficiency compare with Kaolinite especially in Macrofloc group. From the size-class results, specific size ranges of OMAs can be known when influenced by min-

738 eral particles or oil component participating in their flocculation. Since the
739 Bentonite clay is one of the most common mineral particles in natural envi-
740 ronments, its role in absorbing oil, forming OMAs and influencing the fate
741 of oil need to be incorporated in future modeling efforts.

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